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Computation of daily Penman–Monteith reference evapotranspiration in the Carpathian Region and comparison with Thornthwaite estimates

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Abstract. The Pannonian Basin Experiment (PannEx) is a Regional Hydroclimate Project (RHP) of the World Climate Research Programme (WCRP) Global Energy and Water Exchanges (GEWEX) Project. A gridded meteorological dataset is available for the PannEx region as part of the CarpatClim database, which consists of homogenized and harmonized daily meteorological observations for several climate parameters with 0.1° spatial resolution in the period of 1961–2010. The estimation of the Penman–Monteith reference evapotranspiration (ET₀) on the daily scale was performed for the CarpatClim grid as one of the first results in the PannEx initiative. This study compares the already accessible Thornthwaite estimates of potential evapotranspiration (PET_Th) on the monthly scale to the newly derived Penman–Monteith estimates. The comparison is made on an annual and seasonal basis for the 50-year period. The distribution of both estimates is influenced by geographical location and orographic features. The annual time series are similar but the regional-average annual values of ET₀ are ~ 80 mm greater than the Thornthwaite estimate in the whole CarpatClim region. The relative bias and root mean square error was computed as well. The classical Thornthwaite method underestimates the ET₀ by more than 20 % over extensive regions for selected grid points at elevations lower than 200 m in the Pannonian Basin. The slope of the fitted linear trend indicate increasing reference evapotranspiration in the Pannonian Basin due to climate change.

1 Introduction

The Global Energy and Water Exchanges (GEWEX) Hydroclimatology Panel (GHP) provides a framework to assist with the understanding and prediction of hydroclimatic conditions from continental to local scale by improving our knowledge of environmental water and energy exchanges at the regional scale. (GEWEX, 2020 – https://www.gewex.org/, last access: 5 February 2020).

The PannEx is a Regional Hydroclimate Project (RHP) of the GEWEX GHP. It is a GEWEX- initiated, community driven research network in the Pannonian/Carpathian Basin (both names are used). It aims to achieve a better understanding of the Earth System components and their interactions in the Pannonian Basin (PannEx, 2020 – https://sites.google.com/site/projectpannex/, last access: 5 Febru-

ary 2020). The international efforts in PannEx involve research institutions, universities and national meteorological and hydrological services in an integrated approach towards identifying and increasing adaptation capacity to face climate change in the Pannonian Basin.

The PannEx research community has prepared a White book with the objective of first to identify the main scientific issues to address, and then drawing up a detailed science plan to define the major objectives outlined (Ceglar et al., 2018; PannEx White Book, 2019).

The Pannonian Basin (Fig. 1) is enclosed by the Carpathians and the Transylvanian Plateau to the east and north. Over these regions, the Romanian Lowlands and the Subcarpathian depressions beyond the Carpathians (also known



Figure 1. The Pannonian Basin with major orographic and river systems. The area of CarpatClim dataset is shaded.

as Transcarpathia) are in the focus of PannEx initiative (Matenco and Radivojević, 2012; Haas et al., 2013).

The almost-closed structure of the Pannonian Basin in the heart of the European continent makes it a good natural laboratory for studying the water and energy cycles, focusing on the relevant physical processes. However, the data required for estimation of the radiation balance (short- and longwave) and surface energy budget components (net radiation, soil heat flux, sensible and latent heat fluxes) in a daily or finer temporal scale are missing for regional studies. Moreover no long-term coherent soil moisture dataset is available over the Pannonian Basin for studying water and energy cycles. However, the effective precipitation (precipitation - potential evapotranspiration) is an important limiting factor in the agriculture of the Pannonian Basin, and soil moisture content exhibits a strong relationship with dry and wet spells (see also Seneviratne et al., 2006; Torma et al., 2011; Pieczka et al., 2017).

There are various approaches to the definition of the different types of evapotranspiration to describe the evaporative demand of the atmosphere. In the physical sense, actual evapotranspiration (AET) is the sum of the evaporation (E) from the water, soil and plant cover and the amount of water transpired by plants (T). This is often limited by the currently available energy and evaporable water, as well as by characteristics of the plant cover and the soil. Based on these factors, we can distinguish (i) the climatic potential evaporation (C_PET), (ii) the potential evapotranspiration (PET) and (iii) different types of calculation of PET for example the Penman-Montieth reference evapotranspiration (ET_0) or Thornthwaite type potential evapotranspiration (PET_Th) in order to clarify the definitions for general application (McMahon et al., 2013; Lang et al., 2017). Climatic potential evapotranspiration is the maximal possible level of evapotranspiration from a surface saturated in water under given land-use and climatic conditions (Katerji and Rana, 2011). This is the evapotranspiration potential of a given land use. Climatic potential evapotranspiration depends on the characteristics and type of the evaporative surface (bare soil, low and high vegetation, water body, forest, etc.). Thus, climatic potential evapotranspiration is a function of atmospheric forcing and surface types (Rey, 1999; Katerji and Rana, 2011).

The concept of potential evapotranspiration (PET), as described by Penman (1948), is "the amount of water transpired in a unit of time by a short green crop completely shading the ground, of uniform height, and never short of water" (Sellinger, 1996). According to Anon (1956) PET is defined by removing the influence of surface type from the calculation.

On the basis of the Penman–Monteith methodology, reference evapotranspiration (ET_0) was introduced to study the evaporative demand of the atmosphere independent of crop type, crop development and agricultural practices (Allen et al., 1998). ET_0 represents the theoretical evapotranspiration from an assumed reference surface: an extensive surface of green grass of uniform height, actively growing, completely shading the ground, and not short of water.

According to Allen et al. (1998) the use of other denominations, such as potential evapotranspiration, is strongly discouraged due to ambiguities in their definitions. Based on the review papers ET_0 has become one of the most widely used methodologies for calculation of evapotranspiration and evaporative demand (Fisher et al., 2005, 2011; McMahon et al., 2013; Rácz et al., 2013; Lang et al., 2017).

In practice, estimation of actual evapotranspiration (AET) is often made by using information about potential evapotranspiration and (i) soil moisture (e.g. Dyck, 1983; Sellinger, 1996; Ács et al., 2011; Ács, 2017), (ii) plant constants (Allen et al., 1998; Nistor et al., 2016), (iii) standard meteorological data (McMahon et al., 2103) and combined satellite and ground-based measurement (Westerhoff, 2015; Barik et al., 2016). The estimation of evapotranspiration has been widely used not only in agricultural hydrology for planning crop irrigation water needs (Allen et al., 1998), but also in drought monitoring (Vicente-Serrano et al., 2010; Ziese et al., 2014), water balance modelling (Bai et al., 2015), land surface modelling (Niu et al., 2011) and regional biodiversity modelling by ecologists (Currie, 1991), forest hydrology (Rao et al., 2011), and in research of forest ecosystems (Fisher et al., 2005), for example.

Various methods are available to estimate the potential evapotranspiration from standard meteorological observations (Rácz et al., 2013; Lang et al., 2017). The classical Thornthwaite (1948) temperature-based method (PET_Th) is widely used for estimation of evaporative demand, it being calculated only on the basis of air temperature and latitude of the site (Ács et al., 2011). The Penman–Monteith method is considered to be the most physical and reliable method for estimation of potential evapotranspiration and is often used as a standard to verify other empirical methods (Lang et al., 2017).

Our first task in PannEx was the computation of the reference evapotranspiration (ET_0) on a daily scale for the Carpat-Clim grid. The CarpatClim database also provides Thornthwaite estimate of potential evapotranspiration on a monthly basis, which is too coarse for several applications. The objectives of this study are: (i) to estimate (ET_0) for the whole CarpatClim dataset in the period 1961–2010 on the daily scale by using the Penman–Monteith method (Allen et al., 1998; Zotarelli et al., 2010), and (ii) to compare ET_0 with the classical Thornthwaite estimates (PET_Th).

The structure of the applied dataset and methodologies and the first results of the comparisons together with the linear trend values of ET_0 in the period 1961–2010 are presented in the following sections.

2 The CarpatClim dataset

The CarpatClim dataset includes gridded daily minimum, mean and maximum temperature, precipitation, wind direction and speed at a 10 m height, sunshine duration, cloud cover, global radiation, relative humidity, vapour pressure, air pressure and snow depth data along with several climate indicators, 37 in total on a $\sim 10 \text{ km} (0.1^\circ)$ spatial resolution grid. Climatological grids cover the area between latitudes 44 and 50° N, and longitudes 17 and 27° E. The time period of the dataset is 1961–2010. The data can be downloaded from the project website: http://www.carpatclim-eu.org (last access: 5 February 2020; CarpatClim, 2020a) (see also in Szalai and Vogt, 2011; Lakatos et al., 2013).

The target area of CarpatClim partially includes the territories of the Czech Republic, Slovakia, Poland, Ukraine, Romania, Serbia, Croatia, Austria and Hungary. Data from 415 climate stations and 904 precipitation stations were used to create dataset (Lakatos et al., 2013; Spinoni et al., 2015; Cuxart and Matjačić, 2015).

To ensure the use of the largest possible station density, the processing of the homogenization and gridding was implemented by the countries themselves using identical methods and software. The commonly used methods were the MASH (Multiple Analysis of Series for Homogenization; Szentimrey, 1999, 2011) procedure for homogenization, quality control, and completion of the observed daily data series; and the MISH (Meteorological Interpolation based on Surface Homogenized Data Basis; Szentimrey and Bihari, 2007) for the gridding of homogenized daily data series. The uncertainty metrics of the gridded parameters are available as the project deliverable 2.5 (CarpatClim D2.5, 2020b). The harmonization of the data was carried out by exchange of near-border station data between neighbouring countries before and after homogenization. The aim of the harmonization was the matching of national grids so as to remove edge effects. It is noted that snow depth was estimated by ZAMG snow model (Szentimrey, 2012; Olefs et al., 2013). The CarpatClim is widely used for validation purposes for regional climate model simulations and for climate vulnerability assessments (Chervenkov et al., 2019; Torma 2019). Public access to CarpatClim dataset is through the website of the Copernicus Climate Change Service as an external sub-regional dataset (Copernicus CCS, 2020; http://surfobs.climate.copernicus.

eu/dataaccess/access_carpatclim.php, last access: 5 February 2020).

The CarpatClim database also provides Thornthwaite estimate of potential evapotranspiration (PET_Th) on a monthly basis, derived from the geographic position, the average day length (hours) of the month, the monthly average temperature and the monthly and yearly heat indices derived from these variables (Sellinger, 1996; Mihic et al., 2013). The monthly dataset allows the examination of the evaporation conditions of the Pannonian Basin where the monthly time scale is appropriate (Nistor et al., 2016). The CarpatClim drought indices and monthly PET_Th have been used for European drought monitoring and drought assessment (Spinoni et al., 2013) and for evaluation of the EUMETSAT LSA-SAF evapotranspiration products .(Sepulcre-Canto et al, 2014), for example.

3 Daily ET₀ derived by Penman–Monteith reference evapotranspiration method

This study is a substantial step moving from the temperaturebased, monthly Thornthwaite (PET_Th) expression to the fully physical, daily Penman–Monteith ET_0 . Our main applied scientific goal was to develop a procedure for computing of ET_0 on a daily basis for CarpatClim, as the daily values are essential for examining processes at temporal scales for which the monthly Thornthwaite data are insufficient.

The widely used Penman-Monteith Reference Evapotranspiration method, the so-called FAO-56 methodology (Allen et al., 1998) is applicable for the computation of daily ET_0 for the CarpatClim database. The Penman-Monteith method combines both energy and mass balances to model the ET₀. It is based on fundamental physical principles, which guarantee the universal validity of the method. However, it needs a number of meteorological variables which may not be available everywhere. One of the advantages of CarpatClim dataset that all climate variables, required for calculation of ET_0 are available. The method assumes a surface of short grass (0.12 m high) that is well watered, actively growing, completely shades the soil, and has an albedo of 0.23. The necessary variables we used for estimation of the radiation balance and ET₀ are as follows: global radiation (direct and diffuse short wave radiation), temperature (mean, maximum, and minimum), wind speed and relative humidity at 2 m above the ground. The 17-step calculation process published by experts from the University of Florida (Zotarelli et al., 2010) was adapted and coded for the CarpatClim dataset.

The formula used to express reference evapotranspiration $ET_0 \pmod{d^{-1}}$ and which we applied for the gridded daily data is the following Eq. (1):

$$ET_0 = \frac{0.408\Delta (R_n - G) + \gamma \frac{900}{T + 273} u_2 (e_s - e_a)}{\Delta + \gamma (1 + 0.34 u_2)}$$
(1)

where R_n is the net radiation at the crop surface $(MJ m^{-2} d^{-1})$, *G* is the soil heat flux density $(MJ m^{-2} d^{-1})$ (at daily timescales *G* is assumed to be zero), *T* is the mean daily air temperature at 2 m height (°C), u_2 is the wind speed at 2 m height (m s⁻¹), e_s is the saturation vapor pressure (kPa), e_a is the actual vapor pressure (kPa), $e_s - e_a$ is the saturation vapor pressure deficit (kPa), Δ is the slope of the vapor pressure curve (kPa °C⁻¹), γ is the psychrometric constant (kPa °C⁻¹).

4 Comparison of the Penman–Monteith ET₀ and Thornthwaite estimates for CarpatClim dataset

The comparison between Penman–Monteith ET₀ and Thornthwaite PET Th estimates is made on annual and seasonal (spring: MAM, summer: JJA, autumn: SON and winter: DJF) basis. The Fig. 2 shows the annual and seasonal averages (mm) of Penman–Monteith estimates of ET₀ and Thornthwaite estimates for PET_Th over the CarpatClim region for the 50 years from 1961-2010. The spatial distribution of both estimates is influenced by geographical location and the orographic features of the Larger Carpathian region: the Carpathian Chain together with the Carpathian Basin. The highest differences in the annual averages appear at lower elevations in the middle and south-eastern part of the region. The spatial average annual Penman-Monteith estimates are ~ 80 mm greater than the Thorntwaite estimation for the whole CarpatClim region. A similar spatial pattern of the differences can be found in spring, with the Penman-Monteith estimates by 60 mm on a regional average. The summer estimates are in good agreement, although Thorntwaite estimates are lower in the Pannonian Basin, which is the focus area of the PannEx initiative. Areas of differences between the Penman-Monteith ET₀ values and the Thornthwaite estimates appear on the map illustrating autumn. The winter Thornthwaite estimates (PET_Th) are below 20 mm at all the grid points in the CarpatClim region: much lower than the Penman–Monteith estimates which show even 60 mm ET_0 in the lowlands in the southern part of the region.

Further measures of the comparison can be applied such as the relative bias (RE) and the root mean square error (RMSE). As the Penman–Monteith method is considered the most reliable method, the validation was made against that for each grid point on an annual basis. Similar comparisons can be found in Chen et al. (2005). Note that the biases and root mean square errors are normalized by the means of the Penman–Monteith estimates. These two error statistics give an averaged difference over the period 1961–2010. The annual bias indicates that the Thornthwaite method underestimates the potential evapotranspiration in the Carpathian Region (Fig. 3, left panel) relative to Penman–Monteith ET₀. The lowland areas can be characterized by underestimation above 20 % over extended regions, while ~ 4 % overestimation appear in a few gridpoints in the mountainous part of the domain. The right panel in Fig. 3 shows the distribution of annual relative RMSE. The annual relative RMSE ranges from 2.6% to 31.7% across the Carpathian region.

As PannEx focused on the Pannonian Basin we examined the annual cycle of both estimates for selected grid points at elevations lower than 200 m. The monthly averages are shown in the Fig. 4. The Penman–Monteith estimates substantially exceed the Thorthwaite estimates in winter and spring; but from August to October this relationship them reverses. The highest difference is 24 mm in April in favour of Penman–Monteith estimates. The average monthly underestimation from August to October is 4 mm.

Climate change influences all essential climate variables (Spinoni et al., 2015) in the Carpathian region, including the evaporative demand of the atmosphere. The slope of the fitted linear trend to the annual ET_0 is presented in the Fig. 5. Areas with significant trends at $\alpha = 0.1$ significance level are marked with crosses on the map. The increasing trend in the ET_0 is obvious in the Carpathian region from 1961 to 2010. Only small areas (depicted in blue) show decreasing but insignificant trend in these 50 years. The trend of the reference evapotranspiration ET_0 is 0.868 mm yr⁻¹ on average, resulting in a 42.5 mm increase from 1961 to 2010 in the Carpathian region. The highest change is close to 125 mm in flat areas in the north-west part of the region.

5 Concluding remarks and future plans

Penman–Monteith reference evapotranspiration (ET_0) was computed on a daily scale for the CarpatClim grid for the period 1961–2010. One of our main results is that ET₀ values are now available on a daily time step in the CarpatClim database for the applications in the region of the Pannonian Basin Experiment (PannEx). The classical Thornthwaite estimates (PET_Th) were compared on annual and seasonal bases for the 50-year period (1961-2010) to the Penman-Monteith estimates. Both estimates are influenced by geographical location and orographic features. The monthly ET₀ dataset resulted in a more detailed spatial structure and provided higher values, especially in the winter, than the earlier calculated monthly PET_Th in CarpatClim. The annual timeseries are similar, but the spatially averaged annual values of ET_0 are 80 mm higher than the Thornthwaite estimation (PET_Th). The annual relative bias indicates that the Thornthwaite method underestimates the Penman-Monteith estimates in the Carpathian Region, especially in the lowland. As the focus area of PannEx is Pannonian/Carpathian Basin, this is the reason why we examined the annual course of both estimates at elevations lower than 200 m. The highest difference is 24 mm in April in favour of Penman-Monteith estimates. The slope of the fitted linear trend to the annual averages of ET₀ is also presented to illustrate the effect of recent climate change in the Carpathian basin and its sur-



Figure 2. Annual and seasonal averages of ET_0 over the CarpatClim region from Penman–Monteith ET_0 (left panels) and Thornthwaite estimates (PET_Th) (right panels) for the period 1961–2010.



Figure 3. Annual relative bias (%) (a) and RMSE (%) (b) of Thornthwaite estimates (PET_Th) compared with Penman–Monteith ET₀ over the CarpatClim region for the period 1961–2010.



Figure 4. The monthly averages of ET_0 over the CarpatClim region from Penman–Monteith (grey shaded) and Thornthwaite estimates (PET_Th) (black shaded) in the Pannonian Basin at elevations lower than 200 m for the period 1961–2010.



Figure 5. The slope of the fitted linear trend (mm yr⁻¹) to the annual of ET₀ for the period 1961–2010. The significant trends at $\alpha = 0.1$ significance level are marked with stippling.

roundings. We can conclude that both methods are suitable for further investigations in the Carpathian region.

Evapotranspiration is a fundamental component of the hydrological cycle and is essential for understanding landsurface processes in climatology. This was the first step towards creating a realistic daily evaporation database, which is one of the tasks specified in the Science and Implementation plan of the PannEx project (Lakatos et al., 2018). For this purpose, more detailed soil and surface cover databases should be built and reliable soil moisture data (measured or calculated) will be needed (Breuer et al., 2012; Ács et al., 2015). The estimation of the components of the radiation balance will be followed by the parametrization of turbulent fluxes. The standard meteorological measurements of the Hungarian Meteorological Service and high-quality agrometeorological stations (see, for example, the Campbell Scientific products) also provide an opportunity for in situ reference evaporation (LE_0) calculation. Another practical way for calculating reference (ET₀) and/or potential evapotranspiration (based on different methods) is by using operative weather prediction model results (see for example the ME-ANDER nowcasting system based on the high resolution WRF model, Horváth et al., 2015). In practice, estimation of actual ET is often made by using information about potential evapotranspiration and soil moisture. In the future we are considering the development and verification of a unified methodology for derivation of the gridded energy budget components and soil moisture based on standard meteorological measurements and satellite information by using SVAT model approach for the CarpatClim grid system. Moreover, the further extension of the CarpatClim with land-use, soil type, albedo and soil moisture data could support future PannEx activities.

Data availability. Daily Penman–Monteith reference evapotranspiration values based on CarpatClim dataset can be freely download from this link: https://www.met.hu/downloads.php?id=18& file=CARPATCLIM_PM_ET (last access: 19 February 2020).

Author contributions. ML prepared the manuscript with major contributions of TW, further contributions of AH, LH and ZB. ML and TW write the overview of scientific literature and developed the structure of the manuscript together with AH. ZB provided the database and background information of CarpatClim dataset. ML, LH calculated the reference evaporation and made the programming work further contributions of ML and ZB, ML and LH performed the statistical analysis. LH prepared the maps. All authors contributed to the results.

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